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P.J.C. Schrijvers, H.J.J. Jonker, S.R. de Roode, S. Kenjereš

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On the daytime micro-climatic conditions inside an idealized 2D urban canyon

P.J.C. Schrijvers^{a,*}, H.J.J. Jonker^a, S.R de Roode^a, S. Kenjereš^b

^a Faculty of Civil Engineering and Geotechnology, Delft University of Technology, Stevinweg 1, 2628CN Delft, The Netherlands ^b Faculty of Applied Sciences, Delft University of Technology, Van der Maasweg 9, 2629HZ, The Netherlands

Abstract

This study investigated the surface temperature, air temperature and mean radiant temperature inside an idealized 2D street geometry during daytime. The goal was to unravel the relative impact of radiative transfer, heat conduction and ventilation to the urban heat budget. A building-resolving simulation model has been used, which represents these processes at a 1 m spatial resolution. Different combinations of the canyon height to width ratio (H/W) and physical mechanisms were investigated. Shortwave radiation is the main source of energy, and for small H/W can be higher at the canyon ground level compared to flat terrain due to multiple reflections. The longwave trapping effect has the second largest contribution and becomes relatively more important with increasing H/W ratio. The influence of the interior building temperature is small. Surface temperature and mean radiant temperature are closely related, since both are largely controlled by radiative properties. No straightforward relation was found between surface temperature and air temperature, since air temperature is dependent on the competing mechanisms of forced and natural convection. A small increase in air temperature inside the canyon was observed compared to the ambient temperature above roof level. The inclusion of all key physical processes in high detail resulted in large computational requirements. If multiple reflections by the building facades are small, the more traditional, yet much simpler view factor approach will strongly reduce the computational costs as compared to the Monte Carlo technique. The influence of using the view factors on the results must be investigated.

Keywords: urban heat island, mean radiant temperature, urban comfort, surface energy balance, computational fluid dynamics

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1 1. Introduction

During clear nights with weak winds, cities tend to 2 be 1 to 10°C warmer than the surrounding rural envi-3 ronment at night. This phenomenon has been observed for both small and large cities, in the tropics as well 5 as in colder regions, and is called the Urban Heat Is-6 land (UHI) effect [1, 2, 3, 4, 5]. However, the day-7 time UHI effect is much smaller [6, 2] and can even be 8 negative, even in cold climates [7]. Furthermore, the 9 daytime and night time UHI can have different distri-10 bution patterns and intensities over relatively short dis-11 tances of less than 1km, as was found from measure-12 ments by Soltani and Sharifi [8]. This also holds for sur-13 face temperatures and mean radiant temperature, which 14 is a quantity indicative of the human thermal comfort. 15

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The mean radiant temperature only depends on radiation, and is used in the computation of apparent temperatures like the Physiological Equivalent Temperature (PET, Höppe [9]) and the Universal Temperature Climate Index (UTCI, Fiala et al. [10]). Because of the higher temperatures, radiation, and thus heat load, human comfort in the city is more critical during daytime as compared to the night.

Klysik and Fortuniak [2] studied the daytime atmospheric UHI effect of the town Lodz in Poland by using fixed point measurements from a weather station in the city centre over two different periods of three years. They found that on days with clear skies, there are large thermal contrasts within the city. In areas with narrow streets, the air close to the ground may be cooler than the rural environment due to shading of the ground surface. Klysik and Fortuniak [2] state that the radiation and energy budget of roofs play an important role in that scenario. A warm layer of air can be formed at roof level, while the air inside the canyon remains cool due

^{*}Corresponding author

Email address: p.j.c.schrijvers@tudelft.nl(P.J.C. Schrijvers)

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to the shading effect, creating a local inversion layer be-36 tween canyon and roof level and limiting the vertical 37 exchange of air in the street canyons. 38

In addition to air temperature, mean radiant temper-39 ature is also measured in dedicated field campaigns. 40 Such measurements were conducted (amongst others) 41 by Lindberg et al. [11], who observed large local variations of mean radiant temperature. On a large open 43 square in the city centre of Göteborg in Sweden, a 44 peak value in mean radiant temperature of 57°C was 45 found, which is on the threshold for moderate heat stress 46 [12, 13]. In contrast, simultaneous measurements in 47 a courtyard resulted in a mean radiant temperature of 100 48 17°C when shaded, which quickly raised to 27°C when 101 49 the measurement location was directly sunlit. These 102 50 measured values are well below the threshold of mod-51 103 erate heat stress of 55°C, indicating the large impact of 52 104 geometric properties and shading. 53

Even though the mean radiant temperature can be 106 54 measured and the atmospheric UHI frequently ob-107 55 served, the urban UHI is a difficult phenomenon to inter-56 108 pret, due to its spatial inhomogeneity. When interpret-57 ing profiles of thermodynamic variables from a single 58 measurement location, the effect of advection across the 111 59 heterogeneous urban surface should also be taken into 112 60 account, which is difficult to observe or quantify [14]. 113 61 To overcome the local nature of measurements, numeri- 114 62 cal models can be used to study the urban environment, 115 63 in which the complexity and non-linearity of the ur- 116 64 ban environment can be studied in a systematic man-65 117 ner. Often, these models apply on a larger scale (meso-66 scale) and the urban street canyon is parametrized. For 67 119 instance, a meso-scale model (Weather and Research 120 68 Forecasting model, WRF) coupled with a single-layer 121 69 urban canopy model (SLUCM, [15]), was used by 122 70 Ryu and Baik [7]. The building height (H = 15 m) ¹²³ 71 over street width (W = 15 m) ratio (H/W) used was 124 72 73 H/W=1.0. Their study indicated that during daytime 125 the impervious surfaces (including the reduction in sur-126 74 face moisture availability and increased thermal inertia) 127 75 contribute most to the urban heat island $(+2.1^{\circ}C)$. The 128 76 3D urban geometry (transfer of energy in vertical walls, 129 77 shading, radiative trapping and reduction in ventilation) 130 78 actually cools the city (-0.5°C). 79

Ryu and Baik [7] used just a single H/W ratio, though 132 80 it is known that this parameter has a large impact on the 133 81 UHI effect [16]. Marciotto et al. [17] investigated the in-82 fluence of the aspect ratio and mean building height on 135 83 84 local canopy energy fluxes by using an Urban Canopy Model (UCM) similar to Masson [18]. A north-south 85 oriented canyon was used, and a full daily cycle was 86 87

investigated. Results for one time-instance at midday 139 (12:00) were discussed. Increasing the H/W ratio from 0.5 to 10.0, decreases the net radiation, as well as the sensible heat flux, by an amount of 120 Wm⁻² (from 490 to 370) and 300 Wm⁻² (from 360 to 60), respectively. The reduction in absorbed energy is compensated by the conductive heat flux, which transports energy towards the surface, and increases by 180 Wm⁻².

Theeuwes et al. [19] differentiated the two compensating radiative effects in the urban canyon: shadow casting and longwave trapping. The net effect depends on the amount of available shortwave radiation penetrating the canyon. It was found that for H/W=1.0 the largest UHI effect is present. With increasing H/W ratio shading effects start to dominate over the longwave trapping effect during the day, and the UHI is decreasing.

All previously mentioned numerical studies used a meso-scale model in which the urban environment is parametrized. Therefore, results were obtained for the canyon as a whole, or for individual surfaces. When more spatial details are desired, smaller scale models are available that focus on the urban micro-scale. These include for instance Solweig [11], RayMan [20], TUF-3D [21] and envi-MET [22]. Hertel and Schlink [23] developed a method for decomposing the urban heat island intensity at the neighbourhood scale. Envi-MET simulations are used as input, after which the surface energy balance is translated into temperature differences between two neighbourhoods. Unfortunately, no results were presented of a case study, nor validation of the model. In addition to the above models, generic Computational Fluid Dynamics (CFD) models are often applied in which surface temperature is prescribed and air flow is investigated. Toparlar et al. [24] have performed an extensive review of CFD studies on the urban micro climate. They conclude that at present CFD modelling results can be often validated from observations, and CFD models are being increasingly applied for realistic case studies, including the assessment of the effect of adaptation measures. One specific example is Robitu et al. [25], who used a coupled CFD, radiation and conduction model to investigate the effect of vegetation and ponds on the urban micro-climatic conditions. The presence of water ponds and trees improves the urban thermal comfort in summer during day time by cooling the air and shading the urban surface. The influence of trees and ponds was not distinguished, nor the influence of trees and ponds on the different components of the surface energy balance.

Within the current study and [26], a new numerical model (called URBSIM) has been developed, which computes all processes at a 1 m spatial resolution at the

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canyon surface. The goal of the present study was to 189 140 identify the mechanisms that control daytime surface 190 141 temperature, air temperature and mean radiant temper-142 ature within a single canyon at a high resolution. As 191 143 shown in the literature review, these processes are cur-192 144 rently known for the canyon as a whole, thereby neglect-193 145 ing the spatial variability over a single surface. The 194 146 processes taken into account are shortwave radiation, 195 147 longwave radiation, conduction and turbulent transport 196 148 of heat. Our objective was three-fold: 1) disentangle 197 149 the mechanisms involved in the urban heat budget, 2) 198 150 quantify the relation between surface temperature, air 199 151 temperature and mean radiant temperature within a sin-152 gle street canyon, and 3) create insight for future studies 153 into which processes can be parametrized at the urban 200 154 micro-scale. The focus in this study was on the different 2011 155 processes and interactions, not so much on the most ac-156 202 curate representation of the actual urban geometry. This 157 203 was also reflected in our assumption of an idealized 2D 158 204 geometry. 159 205

160 2. Methods and case set-up

The 2D micro-scale model URBSIM, that was dis-161 210 cussed in Schrijvers et al. [26, 27] has been used. URB-162 211 SIM couples a Monte-Carlo radiation model, 1D heat 212 163 conduction equation for the conductive heat flux into 213 164 buildings and the ground and a Computational Fluid 214 165 Dynamics (CFD) model for the convective heat fluxes. 215 166 In the present study, the 2D micro-scale model was ex-167 tended with a new boundary condition for the interior 168 217 building temperature, and mean radiant temperature can 218 169 be routinely computed at any time and at any location. 219 170 A 2D version of the model was used in order to sim-171 220 plify the geometrical complexity. A similar case set-172 221 up is used as in [26], and a range of H/W ratios were 173 considered (0.0, 0.5, 1.0, 2.0). In addition to differ-174 223 ent canyon aspect ratios, also the inclusion of differ-175 224 ent physical processes wereconsidered. Starting from 225 176 a radiation only case, complexity was added by includ-177 ing the conductive heat flux and sensible heat flux. The 227 178 canyon orientation is north-south, such that building fa-179 cades are east-facing or west-facing. This is a typical 229 180 orientation for 2D studies, and was also used in Schri-18 jvers et al. [26]. A spatial resolution of 1 m was used 182 231 at the building surface, and although a full diurnal cy-183 232 cle was modelled, here we only report the situation for 184 233 a solar zenith angle of 28.9° which corresponds to so-185 lar noon for the Netherlands at June 21. Details of the 234 186 model and validation are discussed in Schrijvers et al. 235 187 [26], and only a brief description will be given here for 236 188

convenience. Extensions of the model for this study are discussed in more detail.

2.1. Radiative transfer

Radiative transfer is computed by the Monte-Carlo model that was developed in Schrijvers et al. [26], in which photon paths are computed for four radiative components: 1) diffuse shortwave radiation from the sky, SW_{dif} , 2) direct shortwave radiation from the sky, SW_{dif} , 3) longwave radiation emitted by the sky, LW_{sky} and 4) longwave radiation emitted by the surface, LW_{out} computed as:

$$LW_{\rm out} = \sigma \epsilon_{\rm s} T_{\rm s}^4 \tag{1}$$

with σ the Stefan-Boltzmann constant in $[Wm^{-2}K^{-4}]$, ϵ_s the emissivity of the surface and T_s the surface temperature in each grid cell in [K].

Note that reflection-events are not addressed separately; a photon emitted as direct radiation will be labelled SW_{dir} after a scattering event at the surface. The only exception is LW_{trap} , which is LW_{out} that is absorbed at another surface. Due to the 2D assumption, the azimuthal angle is not taken into account and only the solar-zenith angle is used to describe the solar position. This means that the solar position is only described in the east-west plane, and that the north-south plane is discarded. The azimuthal angle is only taken into account in the computation of the amount of incoming solar radiation at the top of the domain.

The photon packets trajectory is computed from cell face to cell face until a surface is hit. A fraction of the energy $(1 - \zeta)$ is absorbed at the surface, which is related to the albedo of the surface (shortwave radiation) and emissivity (longwave radiation). Note that radiation does not interact with the air inside the canyon, but only interacts at the surface.

The magnitude of the shortwave radiative flux is based on a parametrization proposed by Skartveit et al. [28] and assuming clear skies. Maximum values for downwelling shortwave radiation and the constant value for LW_{sky} that are used at roof level, are shown in Table 1. Although LW_{sky} has a diurnal variation in reality, this is not taken into account in the current study. Similarly, a constant value is used for inlet air temperature and wind speed. Kirchoff's law is assumed for broadband radiation, indicating that the same value is used for absorption (LW_{sky}) and emission (LW_{out}) of longwave radiation at the surface ($\alpha = \epsilon$).

2.2. Mean radiant temperature (T_{mrt})

The existing model has been extended to diagnose the mean radiant temperature, which is computed by

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Radiation					
Emissivity ϵ	0.95	257			
Albedo α	0.40	25			
Latitude	52° 22' N	259			
Longitude	4° 53' E				
Start day	2012-06-10 00:00	260			
End day	2012-06-18 23:59	26'			
max S W _{dir}	$W_{\rm dir}$ 833.1 Wm ⁻²				
$\max S W_{\text{diff}}$	84.2 Wm^{-2}	263			
<i>LW</i> _{sky}	325 Wm^{-2}	264			
	Heat conduction	26			
λ	$0.72 \text{ Wm}^{-1}\text{K}^{-1}$	266			
ρ	1920 kgm^{-3}	267			
$C_{ m v}$	835 Jkg ⁻¹ K ⁻¹	268			
	Computational Fluid Dynamics	269			
Ta	20°C	270			
U	4.0 ms^{-1}	271			
cell width	1.0 m	27			
		273			

Table 1: Input constants for radiation, heat conduction into the urban 275 material and the CFD model.

$$T_{\rm mrt} = \sqrt[4]{\frac{S_{\rm str}}{\epsilon_{\rm p}\sigma}}$$
(2)

282 where S_{str} is the local mean radiant flux density [Wm⁻²] 237 and ϵ_p the emissivity of the human skin, which is a con-238 stant independent of the application with a value of 0.97. 239 The mean radiant flux density can be regarded as the 240 amount of radiation (both shortwave and longwave) that 241 is absorbed by a person. It is computed following Thors-242 son et al. [29] 243

$$S_{\text{str}} = (1 - \alpha_k) \sum_{n=1}^{6} S W_n F_n + \epsilon_p \sum_{n=1}^{6} L W_n F_n$$
 (3)

where n is the orientation (north, east, south, west, top, ²⁹³ 244 bottom), α_p is the albedo of the human body (with a ²⁹⁴ 245 standard value of 0.3), SW_n the total shortwave radiative ²⁹⁵ 246 flux in $[Wm^{-2}]$, LW_n the total longwave radiative flux in 247 $[Wm^{-2}]$ and F_n a geometric factor representing a stand-248 ing human body. A summation is performed over the 4 249 cardinal points (north, east, south, west), for which the 296 250 geometric factor is set to 0.22 for each direction, while 297 251 the geometric factor is set to 0.06 for radiation enter- 298 252 ing from the top and bottom [11]. Since a 2D setting 299 253 is used in this study, information is missing on the two 300 254 faces that are occupying the sides of the canyon. These 255 301

missing radiative fluxes are taken as the average of the two cardinal points that are available. This can physically be seen as computing mean radiant temperature on a square surrounded by obstacles.

2.3. Mean radiant temperature validation

The computation of T_{mrt} within the Monte-Carlo model is validated against values of the actinic flux from Madronich [30]. In that paper, a derivation of the actinic flux (also called integrated density or flux density) is given and solutions are presented for the irradiance for direct and diffuse shortwave radiation.

Two cases are considered. In the first case direct radiation of 800 Wm⁻² has been emitted with a solar zenith angle of 0 degrees onto a diffuse scattering surface with albedo $\alpha = 1$ (no absorption). For this configuration, the horizontal flux should be half of the incoming energy (400 Wm⁻²). The left panel of Fig. 1 shows the horizontal and vertical flux directions, which show that the horizontal component is 400 Wm⁻² (50%) of the incoming direct radiation. There are small spatial differences due to the Monte-Carlo method, which are around 2 Wm^{-2} . The averaged difference is 0.8 Wm^{-2} for the vertical components (0.1%) and 0.15% for the horizontal components. These differences decrease with increasing number of photons.

In the second simulation setup, diffuse radiation (100 Wm⁻²) is emitted, for which [30] derived that the irradiance is equal in all directions (this only holds for a perfect reflecting surface, with $\alpha = 1$). Results are shown in the right panel of Fig. 1, which shows an uniform distribution of 100 Wm^{-2} with fluctuations of 0.5 Wm^{-2} .

The Monte-Carlo radiation model performs well against the results by [30], and as such will be used in the remainder of this study. The spatially averaged Monte-Carlo results are within 0.15% of Madronich [30].

2.4. Conductive heat flux

The conductive heat transfer is computed using the temperature gradient inside the urban material (building walls and layers beneath the street)

$$G_i = -\lambda \frac{\partial T}{\partial x_i} \tag{4}$$

where λ is the thermal conductivity of the material in $[Wm^{-1}K^{-1}]$ (see Table 1) and x_i the distance into the ground or building surface in [m]. The value of λ used in this study is that of brick, which is close to the thermal conductivity of asphalt and medium to dense concrete. Note that this is a highly simplified representation



Figure 1: Validation of radiative fluxes for two numerical tests against analytical results by Madronich [30]. Results show the horizontal flux (eastward and westward) for direct shortwave flux density for all directions and diffuse shortwave flux density for all direction. Input for the direct shortwave radiation is 800 Wm^{-2} and 100 Wm^{-2} for the diffuse shortwave radiation.

of the building walls, in which normally two layers of 322 302 brick are used with a cavity in between. To simplify the 323 303 problem, the cavity is not taken into account. Further- 324 304 more, no sensitivity study on other building materials is 325 305 conducted, which is left for future work. The temper-326 306 ature profile is computed using the 1D heat conduction 327 307 equation 308 328

$$\frac{\partial T}{\partial t} = k_{\rm d} \frac{\partial^2 T}{\partial x_i^2} \tag{5}$$

where k_d is the thermal diffusivity in [m²s], based on the conductivity λ , density ρ in [kgm⁻³] and specific heat C_v of the ground or obstacle in [Jkg⁻¹K⁻¹].

$$k_{\rm d} = \frac{\lambda}{\rho C_{\rm v}} \tag{6}_{335}$$

Two different boundary conditions are used for the 336 309 building interior. The first is a zero flux boundary con- 337 310 dition at a distance of 1 m into the ground or 0.25 m into 338 311 an obstacle. It was found that diurnal temperature cycle 339 312 does not influence the interior temperature in the ground 340 313 over more than 1 m, which is why this distance was 341 314 used. The distance of 0.25 m for the fixed building tem- 342 315 perature is based on the typical thickness of a building 343 316 wall. Alternatively, a fixed interior temperature is used 344 317 318 of 20°C at the same distance into the ground and obsta- 345 cle. The choice of 20°C is based on a temperature in-346 319 side a building that is comfortable. As a consequence of 347 320 the zero-flux boundary condition, the interior building 348 321

temperature follows from the absorbed radiation, sensible heat flux and conductive heat flux of previous time steps. With a constant, prescribed interior temperature, the energy inside the building is dissipated or generated, which can act as an unlimited source of energy. Physically, this can be seen as using an extremely efficient air-conditioning unit which is able to maintain the prescribed interior temperature. A time step of the global model is used of 6 minutes, except for the conductive heat flux where a time step of 1 second is used. The surface temperature of each grid cell is thus updated every 6 minutes. This will be discussed in more detail in section 2.7.

2.5. Computational Fluid Dynamics (CFD) model

Ventilation effects are computed by an in-house developed CFD model [31, 32, 33, 34, 35]. This model uses the Transient Reynolds-Averaged Navier-Stokes (T-RANS) equation to solve the wind field and air temperature distribution, and includes buoyancy effects.

The unknown Reynolds stresses $\overline{u_i u_j}$ [m²s⁻²] are computed using the $k - \varepsilon$ turbulence model, which relates the turbulent stresses to the turbulent kinetic energy k [m²s⁻²] and dissipation ε [m²s⁻³]. The model coefficients ($C_{\mu}, \sigma_k, \sigma_{\varepsilon}, C_{\varepsilon 1}, C_{\varepsilon 2}$ and $C_{\varepsilon 3}$, see table 2) used are taken from the standard k- ε model, as often used for a wide range of turbulent flows in street canyons ([36], [37], [38]).

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C_{μ}	σ_k	σ_{ε}	$C_{\varepsilon 1}$	$C_{\varepsilon 2}$	$C_{\varepsilon 3}$
0.09	1.0	1.3	1.44	1.92	1.44

Table 2: Model coefficients used in the standard $k - \varepsilon$ turbulence model.

In addition to the velocity field, the temperature field 349 is solved by using the T-RANS equations. The unknown 350 turbulent heat flux $\overline{\theta u_i}$ [ms⁻¹K] is computed using the 351 Simple Gradient Diffusion Hypothesis, where the tur-352 bulent flux is related to the temperature gradient and the 388 353 turbulent viscosity 354

$$-\overline{\theta u_i} = \frac{\nu_{\rm t}}{\Pr_{\rm t}} \frac{\partial T}{\partial x_i}.$$
 (7) ³⁹¹
³⁹²

393 with v_t the eddy viscosity $[m^2s^{-1}]$, and Pr_t the turbu-355 lent Prandtl number (the ratio between the eddy diffu-356 sivity for momentum and heat transfer K_m/K_h), and is 357 396 set to 0.86. Although the value of the turbulent Prandtl 358 397 number depends on stability [39, 40], we set Pr_t to 0.86 359 which is typically used in commercial CFD codes and is 398 360 also in between the range between 1/3 and 1 commonly 361 used in large-eddy simulation models for convective and 400 362 401 stable conditions [41]. 363

The sensible heat flux at each individual surface grid cell is defined as

$$SHF_i = \rho c_p \overline{\theta u_i}$$
 (8) 405

where ρ is the density of air (1.208 kgm⁻³) and c_p the ₄₀₇ 364 specific heat capacity of air (1004 $Jkg^{-1}K^{-1}$). The sen-365 sible heat flux is computed based on the temperature 366 gradient between surface and air in the neighbouring 367 grid cell. The Boussinesq approximation is used, stat- 410 368 ing that density differences can be neglected except for 411 369 the buoyancy term. Buoyancy effects are taken into ac-370 412 count in the computation of the temperature field, and 371 413 are therefore not specifically required in the computa-372 414 tion of the sensible heat flux. 373 415

2.6. Integrated energy balance model 374

All sub-models compute a part of the total surface en- 418 375 ergy balance, which dictates that all fluxes should bal- 419 376 ance: 377

$$S W_{\text{dir}} + S W_{\text{dif}} + L W_{\text{sky}} + L W_{\text{trap}}$$

$$= L W_{\text{out}} + S H F + G$$
(9) 424

where SHF is the sensible heat flux (for SHF > 0 there 426 378 is heating of air) and G the conductive heat flux (for $_{427}$ 379

G > 0 energy is added to the ground). The latent heat flux (evaporation of water) is not taken into account in this study. Multiple reflections of radiation are not addressed separately, but are included in the terms of $S W_{\text{dir}}, S W_{\text{dif}} \text{ and } L W_{\text{sky}}.$

The controlling parameter for the surface fluxes is the surface temperature. A skin layer is assumed to prevent large variations in surface temperature in time

$$\Delta_{\rm skin}\rho_{\rm skin}C_{\rm v,skin}\frac{\partial T_s}{\partial t} = \Gamma \tag{10}$$

with Γ the flux imbalance resulting from the surface energy balance and $\Delta_{skin}\rho_{skin}C_{v,skin}=0.01 \text{ JK}^{-1}\text{m}^{-2}$, which results from a very thin layer Δ_{skin} . This also creates an under-relaxation-factor, which helps to stabilize the simulation.

For a time step, all fluxes in the surface energy balance are computed based on the surface temperature of the previous time step in each building surface grid cell. This can result in a small flux imbalance Γ . Based on the flux imbalance and old surface temperature, the surface temperature for the new time step is computed. Surface temperature is thus not a fixed value, but interactive through all surface fluxes.

The time step of 6 minutes is used for the Monte-Carlo radiation model, CFD model and integrated model, while the conductive heat flux model uses a time step of 1 second. The time step of 6 minutes is based on the movement of the shadow: for the cases considered here, the shadow does not travel more than 1 grid cell per time step. Ideally, a smaller time step would be used for the CFD model, but this would increase computation times significantly.

2.7. Test cases and methodology

To study the effect of different physical processes as a function of the H/W ratio, an idealized 2D urban geometry has been used consisting of 10 obstacles which are spaced W=50 m apart. By using an array of obstacles, a fully developed flow pattern is found in the most downwind street canyons, which indicates that more obstacles would not change the flow in the next canyon. The establishment of a fully developed flow after multiple obstacles has been obtained from a sensitivity analysis, and is confirmed by literature [38, 42]. A uniform grid of 1 m is used. From sensitivity studies, it was found that for a convergence of the modelling results, a canyon should be covered by at least 20 x 20 cells to be grid independent. The total number of cells ranges between 400,000 (for H/W=0.5 up to 1,500,000 for H/W=2.0. All buildings are B=25 m wide, while building height is varied between 0 m (H/W=0.0), 25 m (H/W=0.5), 50

m (H/W=1.0) and 100 m (H/W=2.0) (see Fig. 2 for a 480 428 schematic overview). 429 481

For the radiation modelling, periodic boundaries are 482 430 implicitly used at the domain sides, such that radiation 483 431 can only be absorbed at the building surface or reflected 484 432 towards the sky. The CFD model uses an inflow bound- 485 433 ary condition with a prescribed uniform inlet velocity, 486 434 outflow boundary condition with a zero diffusion flux 487 435 for all variables, symmetry boundary condition at the 488 436 top of the domain which does not allow a vertical veloc- 489 437 ity gradient and no-slip walls at the ground and building 490 438 surfaces. 491 439

In addition to changes in building height, different 492 440 components of the surface energy balance are switched 493 on and off. Instead of performing a full factor separation 494 442 analysis like Ryu and Baik [7] where all combinations 443 of different processes were considered, the simplest case 496 444 with only radiation is used as a starting point. From this 497 445 basic case, complexity has been added by adding the 498 446 conductive heat flux process with a zero-flux boundary 447 condition (case 2) and sensible heat flux (case 3). The 448 499 case including the sensible heat flux and fixed interior 449 temperature is addressed separately. 450 500

Eight consecutive days were considered in the mid-451 dle of June, which is the month where the sun is at its 452 maximum zenith angle in the Netherlands. The sim-453 ulated weather conditions correspond to sunny, cloud 454 free weather. The maximum radiative components men-455 tioned in Table 2 are a few tens of Wm⁻² smaller than 456 the maximum values observed in the Netherlands. By 457 using eight days, quasi-steady state results are obtained 458 that are independent on the initial conditions (the daily 459 cycle is repetitive). For a point in the center of the 460 canyon for H/W=1.0, differences between the last three 461 days in the maximum surface temperature are below 462 1°C, while the maximum conductive heat flux differs 463 by 3 Wm^{-2} . The time step that is used is 6 minutes. 464 The inlet air has a velocity of 4 ms⁻¹ and a temper-465 ature of 20°C and both are constant with height and 466 time. The inlet wind speed of 4 ms⁻¹ is used to avoid 515 467 an urban heat island internal circulation that can de- 516 468 velop in the presence of very weak background winds 517 469 [43]. The wind speed of 4 ms⁻¹ is similar to the study ⁵¹⁸ 470 by Draxler [44], and leads to a wind speed above roof 519 471 level of 2.5 ms⁻¹ at the most downwind canyon, which 520 472 is slightly lower than the 30 year average wind speed in 521 473 the month June at weather station De Bilt in the Nether-522 474 lands(www.klimaatatlas.nl). 475

476 Both for daytime and nighttime, constant values in 524 time and height are used for the inlet wind speed and air 525 477 temperature. The reason for using these strongly ideal-478 ized lateral boundary conditions, that ignore the diurnal 527 479

cycle of wind and temperature, is to increase the understanding of the individual processes such as the radiative forcing on the surface energy balance. By using more realistic, time-dependent input parameters, differentiating the different processes from the input parameters becomes much more difficult, which would limit the insights gained from this study. The initial surface temperature is set to 27°C on all surfaces based on expected average surface temperatures over the complete canyon. In total, 8 diurnal cycles are simulated, such that the choice of the initial surface temperature does not influence the results. Other input parameters are shown in Table 2.

Fig. 2 illustrates how further results (Figure 4-7) are plotted, where all vertical surfaces are scaled to a length of 1. This allows us to compare different H/W ratios in a single plot. For H/W=0, the full domain is plotted (length of 1000m) which is plotted as a scaled distance from 0 to 4.

3. Results

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Even though multiple diurnal cycles were computed, the results presented here only show the results for the last computed day and the time instance where the sun reaches its highest position. The last day has been used, to ensure that the initial conditions do not have an influence on the results, and the diurnal cycle is repetitive. For details on the diurnal cycle of surface temperature, the conductive heat flux and sensible heat flux, please see [26]. The highest position of the sun is chosen, since radiation is strongest at this point in time, and will thus lead to the most clear relation between mean radiant temperature and surface temperature. Note that the sun is not directly overhead the canyon, such that there is shading of the west-facing wall.

3.1. Case with radiation only

The first case considers radiative equilibrium. Since there is no conduction or convection, this situation implies that all absorbed radiation should be emitted through longwave radiation. This case acts as the reference case, from which the effect of including additional process can be determined. Absorbed longwave radiation from the sky is shown in Fig. 3a, and is decreasing with increasing H/W ratio due to the reducing sky-view factor. For a flat terrain (H/W=0.0) fluctuations are seen in the absorbed radiation. This is due to the Monte-Carlo method, where a finite number of photons are emitted. As a result, each grid cell receives a slightly different amount of energy. The fluctuations

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Figure 2: Schematic of the case set-up with changing H/W ratio. Ten buildings are spaced W=50 m apart, while building height (H) is varied. Plotting is done according to the inset in the top right, which also includes the solar position with zenith angle of 28.9°. Building orientation is shown in the top left, surface 1 is denoted as 'west-facing', surface 3 is denoted as 'east-facing'.

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are decreasing when more photons are emitted, but this 558 528 comes at a cost of increasing computation time. These 559 529 fluctuations are also seen when buildings are included. 530 560 Fig. 3b shows the absorbed longwave radiation that 561 531 is emitted by the buildings and the ground surface. This 562 532

longwave trapping effect shows an asymmetric pattern 563 533 that is due to the differential solar radiative heating of 564 534 the canyon surfaces. The sunlit surface is warmer, emits 565 535 more radiation and therefore the longwave trapping ef- 566 536 fect in the corner between the street and sunlit wall is 567 537 higher. 538

569 For H/W=0.5 and H/W=1.0, longwave trapping 539 570 peaks at the corner between the ground and the east-540 571 facing wall. Trapping in the corner between the ground 541 and the east-facing wall is higher compared to the cor-542 ner with the west-facing wall, since surface temperature 543 574 in the east-facing wall corner is higher, and thus there is 544 575 more emitted longwave radiation. With increasing H/W545 576 ratio, there is less variation in the longwave trapping. 546 With higher buildings, longwave radiation progresses 577 547 towards two infinite plates that are facing each other, 548 where there is a uniform distribution. The corners are 578 549 of less importance for these larger H/W ratios. 550

The total absorbed shortwave radiation (direct and 551 diffuse component) is shown in Fig. 3c. The trap-581 552 ping effect due to multiple reflections can be observed 582 553 554 at the ground level, where the absorbed shortwave radiation exceeds that of the flat terrain for H/W=0.5 and 584 555 H/W=1.0. The shadow location on the ground can also 585 556 be observed, where H/W=1.0 shows a larger shaded 586 557

area of the ground surface as compared to H/W=0.5. For H/W=2.0, the street level is completely shaded.

Due to our assumption of radiation as the only transport means of heat, the total of all the absorbed radiative fluxes balances the emitted longwave radiation. This is shown in Fig. 3d, which displays a clear peak in emitted radiation for H/W=0.5 and H/W=1.0 at the lower corner between the ground and the east-facing wall, where emitted longwave radiation exceeds that of the flat plate. The longwave trapping effect and the absorbed shortwave radiation have approximately the same magnitude, while the magnitude of the absorbed longwave radiation emitted by the sky is much smaller. There is an imbalance over the canyon of 60 Wm⁻², which is due to the time stepping algorithm, where the emitted longwave radiation is based on the surface temperature computed in the previous time step. For H/W=2.0, the lowest emitted radiation is in the corner between the street and the east-facing wall, at 500 Wm⁻², with the total absorbed radiation at roof level at 800 Wm⁻².

3.2. Surface temperature

From the radiation only case, more physical processes are added to study their effect on the surface temperature. For the radiation case, the surface temperature of a flat plate becomes very high at about 82°C (see Fig. 4a). Note that this is not a prescribed surface temperature, but its high value follows directly from the balance between the absorbed and emitted radiation. These temperatures are in line with Garratt [45], who suggested



Figure 3: Absorbed longwave radiation emitted by the sky, absorbed radiation due to longwave trapping, total absorbed shortwave radiation and emitted longwave radiation by walls in case of radiation only (case 1). Plotting follows the inset at the top right, except for a flat plate. Note the difference in scaling on the vertical axis for each plot.

that surface temperatures in the vicinity of 90 to 100° C ₆₃₉ 587

may occur for dry darkish soils of low thermal conduc- 640 588

tivity (0.1-0.2 $\text{Wm}^{-1}\text{K}^{-1}$), considering a simplified form 641 589

of the surface energy balance equation, utilizing likely 590

upper values of absorbed shortwave flux (1000 Wm^{-2}) 642 591 and screen air temperature (55°C). 592 643

If the conductive heat flux process is added, energy 644 593 is transferred into the ground, reducing its surface tem-645 594 perature by 20°C. The inclusion of the sensible heat flux 646 595 allows for another energy transfer from the solid surface 647 596 to air which even further reduces the surface tempera- 648 597 ture by 30°C. 649 598

If obstacles are added, the effects of shadow casting 650 599 (lower surface temperatures on the west-facing walls) 651 600 and multiple shortwave reflections (higher surface tem-601 perature on the ground and the east-facing wall com-602 653 pared to the flat terrain) are clearly demonstrated for 654 603 obstacles with H/W=0.5 and H/W=1.0 (Figs. 4b and 655 604 4c, respectively). The impact of multiple reflections is 605 smaller for deeper canyons, where only a portion of the 657 606 east-facing wall is directly illuminated. Surface temper-607 658 atures at the ground level exceed these of the flat terrain, 659 608 due to the longwave trapping effect, which is 750 Wm^{-2} 609 in the corner between the street and the east-facing wall. 661 610 In the absence of conduction and sensible heat fluxes 662 611 this additional radiative energy causes excessively high 663 612 surface temperatures of about 120°C. 613 664

The addition of the conductive heat flux influences 665 614 the sunlit and shaded part of the canyon differently. In 666 615 the shaded areas, the conductive heat flux has only a 667 616 small effect. For H/W=2.0, the conductive heat flux 668 617 has the smallest influence. Note that a zero-flux bound-669 618 ary condition is used deep inside the obstacle and the 670 619 ground, which impact will be discussed later. 620

The inclusion of the sensible heat flux filters out sur-672 621 face temperature differences for H/W=0.5, where the 673 622 temperature differs from 32°C close to the east-facing 674 623 wall to 20°C at the corner between the west-facing wall 675 624 and roof. For H/W=1.0, the highest surface temper- 676 625 ature is located in the corner between the east-facing 677 626 wall and the ground. A sharp decrease in surface tem- 678 627 perature is found at the top corner. This is due to a 679 628 local very high turbulent viscosity (v_t), which is di- 680 629 rectly influencing the sensible heat flux following from 681 630 Equation 6. It is known that the standard $k - \varepsilon$ model 682 631 computes too high values of turbulent kinetic energy 683 632 at stagnation points [46]. The too large value of tur-684 633 bulent viscosity is directly increasing the sensible heat 685 634 635 flux. The Durbin time-scale limiter τ may be applied,

which bounds the turbulent viscosity and therefore also 686 636 the turbulent heat flux [46]. Note that the same effect 687 637 is also present for the other H/W ratios, but to a lesser 688 638

extent. For H/W = 2.0, surface temperature is almost uniformly distributed, with temperature differences between 21 and 26°C.

3.3. Surface fluxes

The surface temperature plots reveal, in a qualitative sense, which processes are more important. To get a more quantitative view, all individual fluxes are plotted in Fig. 5. In this plot, all absorbed radiation entering from the sky $(SW_{dir}, SW_{dif} \text{ and } LW_{sky})$ is combined. Also note that positive values of LW_{out} , G and SHF indicate a cooling tendency of the surface.

For the flat plate (Fig. 5a) the absorbed radiation contributes 900 Wm⁻² to the surface, while 450 Wm⁻² is emitted through LW_{out} . The energy surplus is compensated by the sensible heat flux (375 Wm⁻²) and the conductive heat flux (50 Wm⁻²). At this moment in time, there is a flux imbalance of 25 Wm^{-2} .

When cases including obstacles are considered, absorbed radiation remains the largest contribution in the sunlit areas, although its relative contribution decreases for high H/W ratio, which was also shown in Fig. 3. The incoming shortwave radiation is divided over a larger surface area, which results in lower surface temperatures. The longwave trapping effect remains larger than the conductive or sensible heat flux. Compared to absorbed radiation from the sky, the longwave trapping is increasing for increasing H/W ratio.

The conductive heat flux shows very small contributions at the west-facing wall and the top part of the eastfacing walls. Only at the location where the absorbed radiation peaks, the conductive heat flux is significantly transferring energy into the canyon material. The addition of the sensible heat flux shows a small cooling tendency of the west-facing walls. For H/W=0.5, the sunlit parts of the street and the east-facing wall shows a much larger cooling effect, where energy is extracted from the warmer surface. For H/W=1.0 (Fig. 5c), the sensible heat flux at the ground surface is almost zero, which indicates that the temperature differences between the surface and the adjacent air are very small. The sharp peak in the sensible heat flux is clearly seen at the top corner between the roof and the east-facing wall, which is reducing the emitted longwave radiation directly. For H/W=2.0, sensible heat flux increases from 5 Wm⁻² at the ground level to 250 Wm⁻² at the corner between the roof and the east-facing wall, and is generally larger than the conductive heat flux contribution.

3.4. Air temperature

Air temperature and velocity vectors for Case 3 are shown in Fig. 6. For H/W=0.5, a warmer canyon is



Figure 4: Surface temperature for different H/W ratios (sub plots) and different cases (coloured lines). Different vertical scales are used in each plot to better visualize the temperature differences.



Figure 5: Surface fluxes for different H/W ratios (sub plots) and physical processes (coloured lines) for case 3 (including conductive and sensible heat flux, but excluding latent heat flux). The same scale is used for all plots.

observed as compared to the free stream air tempera-741 689 ture, and one single vortex is present. This is in con-742 690 trast with results by Sini et al. [47], who found a second 743 691 vortex in the lower left corner in the absence of buoy-744 692 ancy forces. The non-uniform heating of the canyon in 745 693 the present study suppresses the formation of the second 694 746 vortex. However, in our model simulations a small sec-747 695 ond vortex becomes present only if the buoyancy force 748 696 is switched off. This result (not shown here) hints at a 749 697 subtle effect of the non-uniform canyon heating on the 750 698 suppression of the second vortex. From the east-facing 751 699 wall, air is heated by the surface and forced towards 752 700 the west-facing wall, where the highest air temperature 753 701 is found. Air temperature inside the canyon is 1.5°C 754 702 higher than ambient air, which is comparable in magni-755 703 tude to measurements from Giovannini et al. [48], who 704 found temperature differences up to 2° C for H/W=0.85. 756 705 For H/W=1.0 (Fig. 6b), a similar pattern of the flow 757

706 structure is observed compared to H/W=0.5, but with 758 707 lower air temperatures. Air is locally heated close to 759 708 the surfaces, with a slightly larger region in the corners 709 760 where the wind speed is low. Warm air is transported 761 710 from the west-facing wall upwards towards roof level, 762 711 and from there dispersed to the centre of the canyon due 763 712 to the free stream air flowing over the top. The tem-764 713 perature distribution within the canyon for H/W=1 is 765 714 remarkably different as compared to the other two ex- 766 715 amples shown for H/W equal to 0.5 and 2, respectively, 767 716 in the sense that only a relatively small layer of warm 768 717 air is found near the ground surface. For H/W=1 the 769 718 relatively warm air is more efficiently transported out of 770 719 the canvon. 720 771

Similar to H/W=0.5, a warmer canyon than the free 772 721 stream air is seen for H/W=2.0, with an average air tem-722 perature of 20.1°C. Air temperature inside the canyon is 774 723 rather uniformly distributed, with slightly higher tem- 775 724 peratures along the east-facing wall. One single vortex 725 776 is seen, but velocity inside the canyon is low. Other 777 726 studies also presented a double vortex when one side 778 727 of the canyon is heated, which is not observed here. 779 728 These studies often apply uniform heating on one ver-780 729 tical wall, with typical temperature differences between 781 730 two surfaces of 5°C or 10°C. In the present study, the 782 731 application of an energy balance model at the solid sur-783 732 faces lead to a non-trivial surface temperature distribu-784 733 tion. Therefore, buoyancy forces are not constant inside 785 734 the canyon, which can alter the formation of a double 786 735 or single vortex structure. Further study is required to 787 736 737 fully grasp the influence of these non-uniform tempera-788 ture distributions and when the vortex structure changes 738 inside the canyon. An extensive analysis of air flow in 790 739 deep canyon (H/W=2.1) is provided in the study by Of-791 740

ferle et al. [49]. Measurements were conducted over a range of seasons and primarily analysed for sunny days. A distinction was made on warmer windward and leeward walls in combination with wind directions. When the leeward wall is heated, heat transfer is concentrated near the wall, resulting in vertical transport of heat and less mixing. When this buoyant flow encounters the cross-canyon flow and the shear layer at the canyon top, the different flow layers become well mixed, such that this buoyant flow will be recirculated. Even though there is no cross flow in the present study, the mixing of the shear layer at the canyon top is also observed here. Note that the measurements by Offerle et al. [49] indicate a weaker influence of buoyancy effects compared to several reported numerical studies.

3.5. Influence of interior boundary condition

The zero-flux boundary condition implies that the interior temperature will be determined by the net radiation and the sensible heat flux. To investigate the impact of a zero-flux boundary condition on the results, we have performed a sequence of simulations for H/W up to 1.0 with another boundary condition in which a fixed interior temperature of 20 °C is prescribed.

The surface temperature is plotted in Fig. 7, and compares the Case 3 (including sensible heat flux, but with a zero-flux boundary condition) to the same case but with a fixed T_g . For H/W=0.0, the difference between the two boundary conditions is about 1°C, where the fixed T_g provides a lower surface temperature. For H/W=0.5, there is only a small difference in the surface temperature of about 0.5°C on the east-facing wall resulting from the use of the two different boundary conditions. For all other surfaces, there is no distinctive surface temperature difference when changing the interior boundary condition. For H/W=1.0, the case with the fixed interior temperature provides a higher surface temperature, by 2°C. Furthermore, the decrease in the surface temperature towards the top of the east-facing wall is smaller. The fixed interior temperature is able to extract more energy from the surface into the urban material, compared to the zero-flux boundary condition.

In general, the differences between the two boundary conditions (fixed interior temperature and zero flux) are small, provided that a reasonable estimate is given for the interior temperature. Fig. 8 compares the individual surface fluxes for H/W=1.0. Differences are below 10 Wm⁻² for the conductive heat flux, except for the corner between the east-facing wall and the roof.

Despite the small differences observed in the surface temperatures and fluxes in this study, the fixed interior temperature can act as an unlimited source of en-





Figure 7: Surface temperature for different H/W ratios (sub plots) for case 3 (including conductive and sensible heat flux, with a zero flux boundary condition) and the same case but with a fixed interior temperature T_g of 20°C.

Figure 6: Absolute air temperature (°C) and velocity vectors for Case 3 for different H/W ratios. The x-axis presents the distance from the west-facing wall in meters. Note the different length scales on the horizontal and vertical. A reference vector of 1ms^{-1} is shown on the bottom left of panel a.

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ergy, forcing surface temperature towards the prescribed 842 792

building interior temperature. The large impact of the 793 843 indoor temperature on the outdoor environment has also 794 been demonstrated by Theeuwes et al. [19] who stud-844 795 ied the UHI with the WRF model, and showed a range 845 796 of the UHI between 2 and 7°C when changing building 846 797 847

temperature from 5 to 23°C. 798

The interior temperature boundary conditions there-799 fore can be used as a 'tuning parameter' in which one 849 800 can steer the model results by prescribing the desired 850 801 ground temperature. For instance, when in the valida-851 802 tion case the surface temperature is too low, one can pre-852 803 scribe a higher interior boundary temperature to 'solve' 853 804 the problem. By using a zero-flux boundary condition, ⁸⁵⁴ 805 there is no user input, and the results cannot be con-806 trolled by the user to obtain the desired solution. There-807 fore, the choice was made to use the zero-flux boundary 808 condition as the default option in this study. 809

3.6. Mean radiant temperature 810

The mean radiant temperature at a height of 2 m is 862 811 shown in Fig. 9. For the flat plate (H/W=0.0), a mean 863 812 radiant temperature of 69°C is found. The threshold for 864 813 moderate heat stress is at 55°C, while strong heat stress 865 814 is experienced above 60°C [12, 13], indicating that very 866 815 strong heat stress is experienced. Inclusion of obsta-867 816 cles below H/W=1.0 results in large spatial variations 868 817 due to shading in this case. The mean radiant temper- 869 818 ature in the sunlit part exceeds that of the flat terrain 870 819 due to multiple reflections and longwave trapping and 871 820 peaks at 90°C for H/W=1.0 (extreme heat stress), while 872 821 value of 45°C is found in the shade (no thermal heat а 822 stress). For H/W=0.5, a lower mean radiant tempera-874 823 ture in comparison to H/W=1.0 is found, but extreme ⁸⁷⁵ 824 heat stress is observed over a larger region due to a 876 825 smaller shaded area. For H/W=2.0, there is no sunlit 877 826 part of the canyon at this height, and the mean radiant 878 827 temperature is much more uniform and around 30°C. 879 828 The highest temperature is found near the west-facing 880 829 wall, due to solar reflections from the east-facing wall. 881 830 For H/W=2.0, a nearly constant surface tempera-831 ture at the ground level is found. Surface temperatures 883 832 for H/W=1.0 are higher at the ground level compared ⁸⁸⁴ 833 to H/W=0.5, especially in the sunlit area. The same 885 834 is also seen in the mean radiant temperature. How-886 835 ever, the mean radiant temperature is not only a func-887 836 tion of surface temperature, but also of sky-based radi-837 838 ation (both shortwave and longwave). There is a thus a clear relation between the change in surface tempera-889 839 ture and mean radiant temperature, although linking ab-890 840 solute temperatures show a less trivial relation. 891 841

4. Discussion

4.1. Results

Results in this study showed that for all H/W ratios the surface energy budget is strongly controlled by radiative terms. With increasing H/W ratio, the longwave trapping effect and conductive heat flux become relatively more important. This is similar to the findings by Marciotto et al. [17], who found the largest contribution from the net radiation (all absorbed and emitted radiative terms combined) for H/W=0.5, while for H/W=10.0 the conductive energy flux becomes almost equally important compared to the net radiation.

Mean radiant temperature showed large spatial changes due to shadow locations, where mean radiant temperature is up to 10°C higher compared to air temperature in shaded areas and up to 70°C higher in sunlit areas. Similar results were found by Ali-Toudert and Mayer [50], although the mean radiant temperature in the sunlit part in their study is about 10°C lower than in our study.

The large impact of radiation on the urban energy budget, in particular for low H/W ratios, allows for a relatively simple link between mean radiant temperature and the surface temperature. The strong relation between surface temperature and mean radiant temperature is exploited in the modelling approach followed by simulation models Solweig [11] and RayMan [20], in which the computation of the surface temperature is parametrized but accurate results are obtained for mean radiant temperatures. This study indicates that for H/W=2.0, the sensible heat transfer becomes more important compared to the conductive heat flux, which is highly parametrized in both Solweig and RayMan. This might compromise the accuracy of these models for deep canyons.

The addition of a CFD model in URBSIM gives insight in the impact of surface temperature on the air temperature. There is a complex interplay between forced convection (free stream air) and natural convection. Different vortex dynamics inside a canyon as a function of surface temperature have also been found by Magnusson et al. [51] who found two vortices for a canyon with H/W=2.5. When the leeward wall of the canyon was heated, the weak lower vortex disappeared and one vortex remained. However, these simulations where conducted with a fixed surface temperature, without feedback from warmer/colder air on the surface.

4.2. Model deficiencies

URBSIM has been developed with the goal to reduce the amount of assumptions that are made for modelling



Figure 8: Surface fluxes for H/W=1.0 for case 3 (including conductive and sensible heat flux, with a zero flux boundary condition) and the same case but with a fixed interior temperature T_{φ} of 20°C.

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918 Figure 9: Profiles of mean radiant temperature at z=2m height between buildings 8 and 9 for case 3. 919

921 the urban micro-scale. Therefore, a Monte-Carlo radi-892 ation model and CFD model were used, which repre-893 923 sent the physical processes as detailed as possible. The 894 model results show a small increase in the air temper-895 925 ature inside the canyon for H/W=0.5 and H/W=2.0 as 896 compared to the ambient temperature above roof level. 897 These results appear physically sound. However, a 898 weakness of the model is the representation of the dy-899 929 namics near roof corners, where the CFD model pro-900 duces a too large value for the turbulent viscosity [46], 901 931 which in turn directly affects the turbulent heat flux and 902 reduces the surface temperature locally. 903

In order to take all processes into account in such high 933 904 detail, the urban geometry is highly simplified by only 934 905 taking into account a 2D street canyon. Taking into 935 906

account a full 3D urban geometry would be computationally too expensive. Within URBSIM, it is assumed that if building arrays are much longer than their height, they can be considered to be 2D. Santamouris et al. [52], Coronel and lvarez [53] and Giovannini et al. [48] performed measurements in canyons with ratios L/H of the building length L to the height H in the range between 1.25 to 5.5, and found significant 3D effects on the air flow. All of these measurements observed that flow characteristics inside the canyon are either due to circulatory vortices or finite canyon length effects that are related to 3D flow patterns. These findings indicate that even with very long canyons, 3D effects are present, which are absent in the 2D model used in this study.

In addition to the 2D assumption, there is the effect of temporal variations. Although the whole diurnal cycle is computed, the current paper only focussed on the situation at one time instance. Especially the conductive heat flux shows large diurnal variations. Results shown here are therefore always a function of the situation of the previous time-steps. The large diurnal variations have been reduced in this paper by considering highly idealized conditions, such as a symmetrical 2D street canyon, constant inlet wind speed and air temperature, and a fully developed flow in the canyon of consideration.

Despite these assumptions, calculations are expensive and require substantial amounts of computation time. Similar observation was made by Robitu et al.

[25]. Therefore, the step to a 3D URBSIM program 986 936 is not yet possible, and model simplifications must 987 937 be made. The two most expensive computations are 988 938 the sensible heat flux (CFD model) and the radiation 989 939 (Monte-Carlo method). For the sensible heat flux com-940 putation, no suitable alternatives are present at the mo-941 991 ment, that compute the airflow inside the canyon and handle the complex balance between forced convection 993 943 and buoyancy forces. For the radiation model, a simpler 944 994 model is available using view factor algebra. Especially 995 945 for these simple urban canyon configurations, computa-946 tion times for the radiative components using the view 997 947 factor approach should be small, providing very similar 998 948 results to the Monte-Carlo method. One limitation us- 999 ing view factor algebra is that only a limited number of 1000 950 reflections are typically taken into account (one or two 1001 951 reflections). Therefore, the radiation scheme is the most 1002 952 likely candidate to reduce the computation time, with- 1003 953 out deteriorating the results. 954 1004

One process that has not been investigated in this 1005 955 study is the latent heat flux, which is extracting energy 1006 956 from the urban canyon through evaporation of water. 1007 957 Taking into account the latent heat flux requires one ad- 1008 958 ditional parameter in the CFD model (humidity) and a 1009 959 water balance on the building surfaces. As Theeuwes 1010 960 et al. [54] demonstrated with the aid of a diagnostic 1011 961 equation, the inclusion of the vegetation fraction (di-1012 962 rectly relating to the latent heat flux) is required to ac- 1013 963 curately compute the city-scale UHI effect. 964 1014

965 5. Conclusions

This study focussed on the daytime micro-climatic 1018 966 conditions inside an idealized 2D canyon. A building-967 resolving simulation model has been used, which in-968 cludes key physical processes like radiation, conduction 1019 969 and ventilation by air flow at a 1 m spatial resolution. A range of canyon height to width (H/W) ratios and phys- 1020 971 ical processes are considered. 1021 972 Results showed that the daytime energy budget is 1022 973 strongly controlled by radiation, where absorbed radi- 1023 974 ation from the sky (SW_{dir} , SW_{dif} and LW_{sky} , including ¹⁰²⁴ 975 their multiple reflections) is the main source of energy 1025 976 at the surface, followed by trapped longwave radiation 1026 977 (energy emitted from the surface and absorbed at an- 1027 97 other location). The radiative components are, however, $^{\rm 1028}$ 979 decreasing with increasing building height, while the 1030 980 conductive heat flux is increasing. Mean radiant tem- 1031 981 982 perature increased locally for H/W=0.5 and $H/W=1.0^{1032}$ compared to a flat plate. This is due to increased mul- 1033 983

tiple reflections of shortwave radiation and longwave trapping. For deeper canyons, there is no direct sunlight 1036

reaching the street level, and mean radiant temperature drops quickly.

The link between surface temperature and mean radiant temperature can be made (at least from a quantitative point of view) relatively easy, since both are largely dependent on radiative fluxes. This relation allows for simplified models, where the computation of the surface temperature is parametrized but accurate results are obtained for mean radiant temperatures [11, 20]. The link between surface and air temperature is much harder to make. Air temperature inside the canyon is determined by a complex competition between forced convection (which is a function of free stream wind speed) and natural convection (buoyancy forces). This makes it difficult to develop a parametrization that holds for all H/Wratios and on every individual canyon surface.

The model used in the present study takes all processes into account in high detail, where the input of the user has been minimised as much as possible. The down-side of this approach is that calculations are costly and time-consuming, and the same approach can't be extended to 3D environments due to the large computational requirements of the CFD model and the Monte-Carlo model, in which the number of emitted photons would increase drastically. For these simple building shapes, the Monte-Carlo model may be replaced by a much faster view-factor model.

Furthermore, coupling to a large scale model, such as presented in the overview paper by Chen et al. [55] could be pursued to allow for more detailed meteorological input conditions. Instead of using constant parameters as done now, a diurnal cycle in air temperature and wind speed could be used.

6. Acknowledgment

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- Idealized 2D street canyon simulations are performed computing surface temperature, air temperature and mean radiant temperature
- All key physical processes are taken into account in high detail
- Radiation is the main source of energy to the urban canyon
- No straightforward relation was found between surface temperature and air temperature

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